



Formation Pressure

By:Saeed Fakhr





Pressure

Pressure is defined as force per unit area:



$$P(bar) = \frac{F(dyn)}{A(cm^2)} \qquad P(Pa) = \frac{F(N)}{A(m^2)}$$

 $1 \text{ bar} = 10^5 \text{ Pa} = 14.50377 \text{ psi}$

Density and Specific Gravity



The **density** of a substance is its mass per unit volume:

$$ho = rac{m}{V}$$

$$1ppg = 7.48068 \frac{pcf}{pcf} = 119.83 \frac{kg}{m^3}$$

 $ppg = Ib/gal$

 $pcf = Ib/in.^3$

Specific gravity is the ratio of the density of a substance to the density of a reference substance. The reference substance is nearly always water at its densest (^{ϕ}°C)

$$SG = \frac{\rho}{\rho_{water}}$$

$$\rho_{water} = 1 \ kg/m^3 \rho(kg/m^3) = SG$$

$$\rho_{water} = 8.32 \ ppg\rho(ppg) = SG \times 8.32$$

Hydrostatic Pressure



The pressure at a given depth in a static liquid is a result the weight of the liquid acting on a unit area at that depth plus any pressure acting on the surface of the liquid

$$P_{HYD} = \rho \times g \times h$$

• The pressure at a given depth is independent of direction .it is the same in all directions





• The pressure at a given depth does not depend upon the shape of the vessel containing the liquid or the amount of liquid in the vessel.





Hydrostatic Pressure In Drilling wells

Hydrostatic pressure (P_{HYD}) is the pressure caused by the density or Mud Weight (MW) and True Vertical Depth (TDV) of a column of fluid. The hole size and shape of the fluid column have no effect on hydrostatic pressure.

$$P_{HYD} = 0.052 \times MW \times TVD$$

 P_{HYD} = Hydrostatic Pressure (psi) MW =mud Density (ppg) TVD =true vertical depth(ft) MD= measured depth(ft)



Weight Material



Weighting materials (densifiers) are compounds that are dissolved or suspended in drilling fluid to increase its density. Any substance that is denser than water and that does not adversely affect other properties of the drilling fluid can be used as a weighting material.

Suspend Solid weighting material

Material	Principle Component	Special Gravity
Galena	PbS	Ŷ/Ŷ
Hematite	Fe_2O_3	۵
Barite	BaSO ₃	4/1
Limestone	CaCO ₃	۲/۷
Bentonite		۲/۶

Soluble Salts



Clear Brines Maximum Density





NaCl		NaCl per finished	Percent Volume			
(% by weight)	Density (ppg)	barrel (ppb)	to Start	Cl⁻ (mg/l)	Na+ (mg/l)	NaCl (mg/l)
1	8.39	3.54	99.53%	6,098	3,954	10,052
2	8.45	7.07	99.22%	12,267	7,954	20,221
3	8.51	10.71	98.90%	18,542	12,024	30,565
4	8.57	14.39	98.57%	24,923	16,162	41,085
5	8.63	18.1	98.23%	31,376	20,346	51,722
6	8.69	21.98	97.88%	37,899	24,576	62,475
7	8.76	25.7	97.52%	44,529	28,875	73,404
8	8.82	29.58	97.14%	51,229	33,220	84,449
9	8.88	33.5	96.76%	58,036	37,634	95,670
10	8.94	37.49	96.36%	64,949	42,117	107,067
12	9.07	45.61	95.54%	79,024	51,244	130,268
14	9.19	53.95	94.67%	93,489	60,624	154,113
16	9.32	62.52	93.76%	108,343	70,257	178,600
18	9.45	71.31	92.81%	123,588	80,142	203,730
20	9.58	80.38	91.82%	139,258	90,304	229,562
22	9.72	89.65	90.79%	155,354	100,741	256,095
24	9.86	99.17	89.71%	171,875	111,454	283,329
26	10	108.98	88.59%	188,821	122,444	311,265

DATA FOR SODIUM CHLORIDE SOLUTION



DATA FOR POTASSIUM CHLORIDE SOLUTION

KCI		NaCl per finished	Percent Volume			
(% by weight)	Density (ppg)	barrel (ppb)	to Start	Cl⁻ (mg/l)	K+ (mg/l)	KCl (mg/l)
1	8.39	3.5	99.46%	4,786	5,278	10,064
2	8.44	7.07	99.08%	9,608	10,596	20,203
3	8.5	10.67	98.69%	14,500	15,991	30,491
4	8.55	14.35	98.29%	19,464	21,465	40,928
5	8.6	18.03	97.88%	24,498	27,017	51,515
6	8.66	21.77	97.47%	29,603	32,647	62,250
7	8.71	25.55	97.04%	34,744	38,316	73,060
8	8.77	29.4	96.60%	39,955	44,064	84,019
9	8.82	33.29	96.16%	45,238	49,889	95,127
10	8.88	37.21	95.70%	50,556	55,754	106,310
12	8.99	45.23	94.76%	61,440	67,757	129,197
14	9.11	53.46	93.78%	72,607	80,073	152,680
16	9.22	61.86	92.77%	84,023	92,663	176,686
18	9.34	70.47	91.71%	95,758	105,605	201,362
20	9.46	79.33	90.62%	107,741	118,820	226,560
24	9.7	97.63	88.33%	120,043	132,387	252,430



Material Balance

The concept of a material balance is based on the law of conservation of mass that states that mass can be neither created nor destroyed. Simply stated, the sum of the components must equal the sum of the products. This concept is valid for mass and atoms, but it is not always valid for solutions and compounds due to solubilities and chemical reactions. Mathematically, the concept of the material balance is divided into two parts:

• The total volume equals the sum of the volumes of the individual components.

$$V_{total} = V_1 + V_2 + V_3 + V_4 + \cdots$$

• The total mass equals the sum of the masses of the individual components.

$$m_{total} = m_1 + m_2 + m_3 + m_4 + \cdots$$

$$V_{total}\rho_{total} = V_1\rho_1 + V_2\rho_2 + V_3\rho_3 + V_4\rho_4 + \cdots$$

Pore Pressure



During a period of erosion and sedimentation, grains of sediment are continuously building up on top of each other, generally in a water filled environment. As the thickness of the layer of sediment increases, the grains of the sediment are packed closer together, and some of the water is expelled from the pore spaces. However, if the pore throats through the sediment are interconnecting all the way to surface the pressure of the fluid at any depth in the sediment will be same as that which would be found in a simple colom of fluid. The pressure in the fluid in the pores of the sediment will only be dependent on the density of the fluid in the pore space and the depth of the pressure measurement (equal to the height of the colom of liquid). it will be independent of the pore size or pore throat geometry.

$$P_P = 0.052 \times \rho_P \times TVD$$

P_P=Formation Pore Pressure,(psi)
 ρ_P=Density of Pore Fluid ,(ppg)
 TVD =True Vertical Depth from Mean See Level ,(ft)

Pore Pressure Gradient



Pore pressure gradient is the Pore pressure increase per unit of vertical depth.

The datum which is generally used during drilling operations is the drill-floor elevation but a more general datum level, used almost universally, is <u>Mean Sea Level, MSL</u>. When the pore throats through the sediment are interconnecting, the pressure of the fluid at any depth in the sediment will be same as that which would be found in a simple column of fluid and therefore the pore pressure gradient is a straight line.



 P_{PG} = Pore Pressure Gradient,(psi/ft)) ρ_P =Density of Pore Fluid ,(ppg)





Fluid	Density (lb/gal)	Pressure Gradient (psi/ft)
Freshwater	8.335	0.433
Seawater	8.55	0.444
Saltwater (100,000 ppm)	8.95	0.465
Saturated saltwater (10 lb/gal)	10.0	0.520
16-lb/gal mud	16.0	0.832

Area	Average Pore Fluid Density (lb/gal)	Pore-Pressure Gradient (psi/ft)
West Texas	8.33	0.433
Gulf of Mexico	8.95	0.465
North Sea	8.70	0.452
Malaysia	8.50	0.442
Mackenzie Delta	8.50	0.442
West Africa	8.50	0.442
Anacardo Basin	8.33	0.433
Rocky Mountains	8.39	0.436
California	8.45	0.439



The percentage of pore volume or void space, or that volume within rock that can contain fluids. Porosity can be a relic of deposition (**primary porosity**, such as space between grains that were not compacted together completely) or can develop through alteration of the rock (**secondary porosity**, such as when feldspar grains or fossils are preferentially dissolved from sandstones). Porosity can be generated by the development of fractures, in which case it is called **fracture porosity**.

$$\emptyset = \frac{V_v}{V} = \frac{\rho_m - \rho_b}{\rho_m - \rho_f}$$

 $\emptyset = \mathsf{Porosity}$

$$V =$$
Bulk Volume

Porosity

 V_v =Pore volume ρ_m =Matrix density ρ_b =Bulk Density

 ρ_f =Pore fluid density

 $1 > \emptyset > 0$



Permeability



The ability, or measurement of a rock's ability, to transmit fluids, typically measured in darcies or millidarcies. The term was basically defined by Henry Darcy, who showed that the common mathematics of heat transfer could be modified to adequately describe fluid flow in porous media. Formations that transmit fluids readily, such as sandstones, are described as permeable and tend to have many large, well-connected pores. Impermeable formations, such as shales and siltstones, tend to be finer grained or of a mixed grain size, with smaller, fewer, or less interconnected pores.

$$\boldsymbol{k} = \boldsymbol{A}\left(\frac{\boldsymbol{Q}}{\boldsymbol{\mu}}\right)\left(\frac{\Delta \boldsymbol{P}}{\boldsymbol{L}}\right)$$

k =Permeability *Q*=Rate of flow
μ=Fluid viscosity









Oil and Gas are typically found in SEDIMENTARY ROCKS.



SHALE: has 'platy' particles, like decks of cards. There is some porosity but very little permeability. Until recently shales were considered to be source rocks. Many shales are now considered as "unconventional reservoirs" and require fracturing (or 'fraccing') to induce oil and gas to flow.

SANDSTONE: has 'spherical' particles and typically has good porosity and good permeability. The 'best' reservoirs are frequently sandstone reservoirs.

LIMESTONE: has irregular shaped particles like the shells that they are made of. The porosity is typically 'vuggy' or irregular as are the connections between the pore spaces. Limestone porosity and permeability is not as easy to quantify or predict as that of sandstones but they still can be very good reservoirs.

Overburden Pressure



The vertical pressure at any point in the earth is known as the overburden pressure. The overburden pressure is defined as the pressure exerted by the total weight of overlyingformations above the point of interest. The total weight is the combined weight of both theformation solids (rock matrix) and formation fluids in the pore space.

$P_0 = 0.052 \times \rho_b \times TVD$

$$P_{OG} = 0.052 \times \rho_b$$

 σ_{ob} =Overburden pressure(psi)

 ho_b =Formation bulk density (sediment plus fluid)(Ib/gal)

 σ_{og} =Overburden gradient(psi/ft)

TVD=True Vertical Depth(ft)



A useful equation for calculating the formation bulk density (density of sediment plus fluid) under field conditions of varying lithological and pore fluid density is given by:

$$\boldsymbol{\rho}_{\boldsymbol{b}} = [\boldsymbol{\rho}_{\boldsymbol{m}} - (\boldsymbol{\rho}_{\boldsymbol{m}} - \boldsymbol{\rho}_{\boldsymbol{P}})\boldsymbol{\emptyset}]$$

 ρ_b =Formation bulk density ρ_m =Matrix Density \emptyset =Porosity (Fractional) ρ_P =formation fluid density

Lithology	Matrix Density (g/cc)
Sandstone	2.65
Limestone	2.71
Dolomite	2.87
Anhydrite	2.98
Halite	2.03
Gypsum	2.35
Clay	~2.7-2.8
Fresh Water	1.0
Salt Water	1.15
Oil	0.80



Total overburden pressure is supported by the rock in two ways. The first is through intergranular pressure σ_z , a **matrix stress** caused by the force transmitted through **grain-to-grain mechanical contact**. As formations are compacted by the overburden with increasing burial depth, pore fluid escapes so that **pore pressure** P_f is equal to the hydrostatic pressure of the pore fluid density.

 $P_0 = P_P + \sigma_z$

 σ_{ob} = Overburden pressure P_f = Pore Pressure

 $\pmb{\sigma}_{\pmb{z}}$ =Matrix stress









$P_0 = P_P + \sigma_z$

 σ_{ob} = Overburden pressure P_f = Pore Pressure σ_z = Matrix stress









Compaction

Compaction is the reduction of the sediment bulk volume and is equivalent to volumetric strain:

$$\epsilon_v = \frac{V}{V_0}$$

I/

V =load bearing volume

 V_0 = unloaded initial volume



Compacting Theory



- Compaction is caused by overburden load, thus sediments are compacted more with increasing burial <u>depth</u>.
- Compaction mainly decreases <u>porosity</u>, but also reduces the <u>grain volume</u>. Generally, the rock and pore volumes are reduced with reversible (elastic) and irreversible (plastic) contributions







• Hydrostatic equilibrium within the compacted layers is retained as long as the expelled water



• If the expelled water is not free to escape, abnormal pressures may result. Sufficient compaction cannot occur so the pore fluids carry more of the overburden

Abnormal Pressure

A **"normally" pressured** formation has a pore pressure equal to the hydrostatic pressure of the pore water. Any deviation from the normal hydrostatic pressure environment is referred to as abnormal. High pressures are called geopressures, overpressures or sur-pressures. Low pressures are called underpressures or subpressures.





Normal Pressure

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A "normally" pressured formation has a pore pressure equal to the hydrostatic pressure of the pore water. Since many more wells are drilled in sediments characterized by $\Lambda/\$ bl/gal saltwater, a "normal" pressure gradient, for the purposes of this discussion, is considered to be $\Lambda/\$ psi/ft.



Abnormal Underpressure



Subnormal (low) pressures are encountered in zones with pore pressures lower than the normal hydrostatic pressure.

- Subnormal pressure conditions often occur when the **surface elevation** of a well is much higher than the subsurface water table or sea level.
- Another common cause of abnormally low pressures is depleted sands. These are sands whose original pressure has been depleted or drained away. Depleted sands are found most frequently in reservoirs from which oil and gas have been produced, a common phenomenon in many so-called "mature" oil and gas areas.





Abnormal Overpressure

Abnormal Overpressure characterize those zones that have pore pressures higher than the normal hydrostatic pressure of the pore fluids .Abnormal Overpressure are created and maintained by the restriction of pore fluid movement. Abnormal overpressures are always caused by a particular zone becoming "**sealed**" or isolated. Seals are impermeable layers and boundary zones that will not permit the release of pressure generated by the percolation of fluids and gases to higher zones and subsequently to the surface.

The zone between the normally pressured zone and the overpressured zone is known as the **transition zone**. The pressures in both the transition and overpressured zone is quite clearly above the hydrostatic pressure gradient line. The transition zone is therefore the seal or **caprock** on the overpressured formation.





simplified model consisting of a vessel (drainage path open), the increased load must containing a fluid (representing the pore fluid) and a spring (representing the rock matrix). The overburden stress can be simulated by a piston being forced down on the vessel. The overburden (S) is supported by the stress in the spring (σ) and the fluid pressure (p). If the overburden is increased (e.g. due to more sediments being laid down) the extra load must be borne by the matrix and the pore fluid. If the fluid is prevented from leaving the pore space (drainage path closed) the fluid pressure must increase above the hydrostatic value. Such a formation can be described as overpressured (i.e. part of the overburden stress is being supported by the fluid in the pore spaceand not the matrix). Since the water is effectively incompressible the overburden is almost totally supported by the pore fluid and the grain to grain contact stress is not increased. In a

The compaction process can be described by a formation where the fluids are free to move be taken by the matrix, while the fluid pressure remains constant. Under such circumstances the pore pressure can be described as Normal, and is proportional to depth and fluid density.





Origin of Overpressured Formations



\-Undercompaction

In a **normally pressured** environment, sediments are compacted as the increasing weight of the overburden squeezes out the connate water. Thus, the porosity (void space) decreases with depth.

In the most common scenario, a seal has formed, trapping pore water so that compaction (due to increasing overburden with depth) does not occur as in a normally pressured environment. When the sediments are not compacted enough to form grain-to-grain contact, the overburden is supported in part by the pore pressure, causing abnormally high pore pressure Under **abnormal conditions**, the water is not allowed to escape, and the compacting process is altered. The porosity no longer continues to decrease and, In most cases, will increase below the top of the geopressured zone.







One cause of geopressure is the geologic uplifting displacement of a formation, which physically places a higher-pressured formation from a greater depth to a shallower depth. When a previously normal pressure zone at great depth is displaced by tectonic actions to a shallower depth with the seals remaining intact, the resulting pressure gradient will be abnormally high.



***-Salt Diaparism**

Diapirism is the piercement of a formation by a plastic, mobile, less dense underlying formation, typically salt. Salt exhibits plastic behaviour at elevated temperatures and pressures and due to its low density will move upwards to form salt domes in overlying formations. The creation of the salt dome can lead to abnormal pressure development in surrounding formations in two ways: Firstly, the movement of the salt creates additional tectonic stresses within the overlying sediments whilst at the same time providing a lateral seal limiting pore water expulsion. These tectonic stresses usually create folding and faulting in the surrounding zones. Secondly the salt may encapsulate rafters of overlying formations (usually limestones and dolomites) as it flows upwards, trapping pressures within the

rafters. In this instance the salt prohibits further de-watering of the rafter and abnormal pressure will develop.







The diagenetic changes which occur in shales are one of the most important mechanisms by which abnormal pressure may be generated in marine environment. On initial burial. а marine clays are composed of predominantly smectite clays of which montmorillonite is by far the most common. Montmorillonite has a swelling lattice and contains approximately $\vee -\Lambda \Delta \%$ water during initial deposition. The water is held as interlayer water between the clay platelets and also as free pore water. This environment is usually alkaline in nature and is rich in calcium and magnesium ions but poor in potassium ions. Upon further burial, compaction expels most of the free pore water and the water content is thus reduced to approximately ^{\(\nu\)}. With further burial, there will be increases in both the overburden

load and temperature and these two effects cause all but the last layer of structural water to be expelled to the pore space.

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This causes the clay lattice to collapse and in the presence of potassium ions, montmorillonite diagenesis to illite occurs. If the water released in this process cannot escape during compaction, then the pore fluid will support an increased portion of the overburden and will thus be abnormally pressured.





a. MONTMORILLONITE BEFORE DIAGENESIS

b.LOSS OF SOME PORE WATER AND INTERLAYER WATER



d.FINAL STAGE OF COMPACTION



c.LOSS OF LAST INTERLAYER CONVERT MONTMORILLONITE TO ILLITE



A- Diagenesis of Sulphate Formations

The conversion of gypsum ($CaSO_4 \bullet 2H_2O$) to anhydrite ($CaSO_4$), an extremely impermeable evaporate expels water with increasing depth (pressure and temperature) resulting an abnormally pressured sediments below them.

$$CaSO_4 \bullet 2H_2O \xrightarrow{}_{pressure and Heat} CaSO_4 + 2H_2O$$



Formation Fracture Pressure

Formation Fracture Pressure



Fracture pressure is the pressure required to rupture a formation physically, allowing entry of drilling fluids into the formation.

In most oilwell applications, the rock under consideration is subjected to in-situ stresses with have no shear stresses. Normal stresses which have no associated shear stresses are described as Principal Stresses and are mutually perpendicular to each other:

- ${m \sigma}_1$ = The maximum principal stress
- σ_2 = The intermediate principal stress
- σ_3 = The minimum principal stress:

 $\sigma_1 = P_0$

 σ_1 = σ_H

 $\sigma_3 = \sigma_h$





In most cases the maximum principal stress will be vertical due to the pressure of the overlying rock. This is defined as the overburden pressure.

In a tectonically stable basin, the maximum principal stress will be vertical and the horizontal stresses will be equal.

$$\sigma_1 = P_0$$

σ_1 = σ_3

In any area undergoing tectonic activity the horizontal stresses are distorted creating an intermediate and a minimum principal stress. The three stresses are: vertical stress (overburden, P_0) is the maximum stress and the intermediate horizontal stress σ_H and the minimum horizontal stress σ_h is normal to σ_H

$$\sigma_1 = P_0$$

$$\sigma_1 = \sigma_H$$

$$\sigma_3 = \sigma_h$$



In areas of high tectonic activities such as mountain belts, the maximum principal stress may be horizontal and the minimum principal stress is vertical.



Fracture of the formation results when the pressure in the wellbore is equal to, or greater, than the minimum principal stress (assuming the tensile strength of rock is negligible). The fracture will propagate along the path of least resistance which is perpendicular to the direction of the minimum principal stress. Thus fractures will be vertical in areas where the minimum principal stress is horizontal, and horizontal where the minimum principal stress is vertical.





Estimating Fracture Pressure

Hubbert and Willis:

$$P_F = \frac{(P_0 - P_P)}{3} + P_P$$
$$P_{FG} \approx \frac{(P_{0G} - 2P_{PG})}{3}$$

Under normal hydrostatic conditions:

 $P_{OG} \approx 1 \ psi/ft$

 $P_{PG} \approx 0.465 \, psi/ft$

Then:

 $P_{FG} = 0.64 \, psi/ft$

Matthews & Kelly:



$$P_{FG} = P_{PG} + K_i (P_{OG} - P_{PG})$$
$$P_{FG} = P_{PG} + K_i (\sigma_z)$$

K_i =matrix-stress coefficient





Ben Eaton:

$$\boldsymbol{P_{FG}} = (\boldsymbol{P_{OG}} - \boldsymbol{P_{PG}}) \left[\frac{\boldsymbol{v}}{1 - \boldsymbol{v}}\right] + \boldsymbol{P_{PG}}$$

v= Poisson's ratio of the rock





Measuring Fracture Pressure

Many problems exist in trying to estimate fracture pressures. This is because the exact values of the components that contribute to formation strength are not known. These factors are local.Data from one area cannot be readily applied to other areas.

The fracture pressure estimate is used to help design a drilling program for a well with regard to casing depths and hole sizes. Once a well has spudded, the formation fracture pressure should be determined by physical tests.

Two tests are used to measure the formation strength or fracture pressure. These are: () the Leak-Off Test (LOT) and ($^{\gamma}$) the Formation-Integrity Test (FIT). These tests are conducted after casing has been set and the casing shoe has been drilled out. Although procedures differ from one operator to the other, a common practice is to drill either to the first sand or $^{\gamma} \cdot$ ft of new formation before running either test. In some cases, the test can be run again after drilling further. This is usually done when mud weights are used that exceed those planned for in the well plan or mud program.

Leak-off tests and formation-integrity tests are very similar. The difference is that the leak-off test fractures the formation and measures the actual strength of the formation, while the formation integrity test measures the formation to a predetermined pressure but does not cause a fracture. The formation being drilled often determines which of them will be used. The formation integrity test is used more often in hard-rock formations than the leak-off test.

Leak Off Test (LOT)



1. Drill out the casing shoe and sufficient new formation.

7. Circulate the drilling fluid to ensure a consistent mud weight.

 γ . Stop the rig pumps and shut the well in.

*. Pump mud into the shut-in well at a very low rate. A typical pump rate of $\cdot/\uparrow \diamond$ to \cdot/\diamond barrels per minute (bbl/min) is used. Normally, a cementing unit is used so an accurate reading of volume and pressure can be obtained.

▲. Record the pressure and volume pumped. A graphical presentation should be made of these data to determine the point at which the fluid is being pumped into the formation (leak-off). A normal leak-off test will show the pressure increasing in a straight line with the volume of mud pumped. Once the fracture pressure is reached, the pressure will stop increasing with volume pumped as the fracture is being propagated. The pressure may actually decrease as fluid is pumped into the formation. is an illustration of a graph obtained from a leak-off test.





• Once leak-off has been observed ,stop pumping and observe the well. The pressure should remain relatively the same or decrease slightly once pumping has stopped.

 $P_F(psi) = mw(ppg) + \frac{Leak \ Off(psi)}{0.052 \times TVD(ft)}$ 2.0VERBURDEN LEAK- OFF Pressure (psi x 1,000) PARTICULATE 1.5 FLUID 0.0-0-0-0-0-0 1.0Shut-in+ time (min) A = Leak-off pressure 0.5 B = Maximum test pressure C = Minimum formation stress D = Fracture closure pressure HORIZONTAL 0.0 2 3 5 0 1 4 Volume (bbl)

V. Record the pressure where the fluid started leaking off into the formation .
 Convert this pressure to a mud weight equivalent by using the following equation.

Casing Design Process



The selection of casing string and setting depth is based on formation pore pressure and fracture gradient of the well.





The solid lines in the chart are not accounted for safety factor; therefore, for the first step of casing seat design, safety margin must be applied. For this example, we will add \cdot / $^{\circ}$ ppg for safety for both pore pressure and fracture gradient (Figure $^{\circ}$). You need to add the safety factor into formation pressure and subtract it from the fracture gradient. What's more, the safety factor value may depends on where you work and how much confident in your data.



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Top Down Casing Design

This design will start from the surface of the well down to the bottom and the setting depths are designed within the safety factor limits (dotted lines). We start by drawing the vertical line from the facture gradient dashed line (point A) down to pore pressure dashed line (point B). See Figure ^V. The first casing should be set from surface to $r' \cdot \cdot \cdot'$ TVD.





Next, draw the horizontal line from Point B to Point C located in the fracture dashed line curve. Then draw the vertical line from Point C to intersect the formation pressure dashed line curve at Point D. This is the section casing string which should be set from $r' \cdot \cdot \cdot'$ TVD to $r' \cdot \cdot \cdot'$ TVD.





Applying the same concept to the next string, draw the horizontal line from Point D to intersect the fracture gradient with safety factor chart at Point E and draw the vertical line from Point E to the target depth at Point F. The last casing string should be set from $\hat{\gamma}$, \cdots 'TVD to γ , \cdots 'TVD.





Based on the tow down design concept, we will need to have Υ strings of casing set at $\Upsilon' \cdot \cdot \cdot'$ TVD, $\varphi' \cdot \cdot \cdot'$ TVD and $\Upsilon' \cdot \cdot \cdot$ TVD.





Bottom Up Casing Design

This design will start from the bottom of the well up to surface and the setting depths are designed within the safety factor limits (dotted lines). Starting at the bottom (formation pressure dashed line – Point A), draw a vertical line upward to fracture pressure dashed line – Point B (Figure \mathcal{V}). Casing should be set from $\mathcal{V}^{\prime} \mathcal{V} \mathcal{V}$ TVD to $\mathcal{V}^{\prime} \mathcal{V} \mathcal{V}$ TVD because you can reach TD ($\mathcal{V}^{\prime} \mathcal{V} \mathcal{V}$ TVD) with highest equivalent mud weight and you will not break the formation at shallow depth ($\mathcal{V}^{\prime} \mathcal{V} \mathcal{V}$ TVD). We will apply this same concept to another string





The next casing string is determined by drawing a horizontal line from Point B to intersect the pore pressure dashed line at Point C. Then draw a vertical line from Point C to the fracture gradient dashed line at Point D (Figure ⁺). The Casing must be set from $\gamma' \wedge \cdot \cdot'$ TVD to $\gamma' \wedge \cdot \cdot'$ TVD.</sup>





With the same idea, the next casing string is determined by drawing a horizontal line from Point D to Point E and a vertical line from Point E to Point F (Figure $^{\circ}$). The Casing must be set from surface to $^{\prime}A \cdot \cdot'$ TVD.





